Measurements of atmospheric carbon dioxide (CO₂) and oxygen (O₂) concentrations continue to provide strong insights into human-induced deviations in the global carbon cycle. The rise in atmospheric CO₂ has been well documented since 1958, when high-precision measurements began at Mauna Loa, under the direction of C.D. Keeling, as shown in the left-hand figure. The practices applied at that time to ensure high-quality measurements were adopted for observations by Global Atmosphere Watch participants. The rise of CO₂ concentration has been only about half of what would be expected if all the excess CO₂ from the burning of fossil fuels stayed in the air. The other half has been absorbed by the land biosphere and the oceans, but the split between land and oceans is not easily resolved from CO₂ data alone. This is where O₂ measurements prove useful.

One of the longest O₂ time series, from flasks collected at Cape Grim, Tasmania and analysed at the Scripps Institution of Oceanography, is shown in the right-hand figure. Oxygen abundance is reported as relative changes in the oxygen/nitrogen (O₂/N₂) ratio. A decrease of 100 per meg corresponds to the loss of 100 O₂ molecules for every million O₂ molecules in the atmosphere. Atmospheric O₂ has decreased very slightly, as it is consumed during combustion of fossil fuels. The observed decrease in O₂ is less than predicted by the amount of fossil fuels combusted, because, as plants take up CO₂ through photosynthesis, O₂ is returned to the atmosphere. The O₂ offset therefore quantifies the magnitude of the land CO₂ sink. The ocean CO₂ sink can then also be calculated based on the requirement that the combined sinks sum to the total established from the CO₂ data. The O₂ data also resolve fluctuations seasonally and on other timescales, which provide additional information about the large-scale functioning of the Earth’s biosphere that complements CO₂ data. (Figures courtesy of R. Keeling, Scripps Institution of Oceanography, USA)

Executive summary

The latest analysis of observations from the WMO Global Atmosphere Watch (GAW) Programme shows that the globally averaged mole fractions of carbon dioxide (CO₂), methane (CH₄), and nitrous oxide (N₂O) reached new highs in 2013, with CO₂ at 396.0±0.1 ppm[^2], CH₄ at 1824±2 ppb[^3] and N₂O at 325.9±0.1 ppb. These values constitute, respectively, 142%, 253% and 121% of pre-industrial (before 1750) levels. The atmospheric increase of CO₂ from 2012 to 2013 was 2.9 ppm, which is the largest year to year change from 1984 to 2013. For N₂O the increase from 2012 to 2013 is smaller
than the one observed from 2011 to 2012 but comparable to the average growth rate over the past 10 years. Atmospheric CH$_4$ continued to increase at a rate similar to the mean rate over the past five years. The National Oceanic and Atmospheric Administration (NOAA) Annual Greenhouse Gas Index shows that from 1990 to 2013 radiative forcing by long-lived greenhouse gases (LLGHGs) increased by 34%, with CO$_2$ accounting for about 80% of this increase. Uptake of anthropogenic CO$_2$ by the ocean led to increased CO$_2$ concentrations and increased acidity levels in seawater. During the last two decades ocean water acidity expressed as pH decreased by 0.0011–0.0024 units per year, and the amount of CO$_2$ dissolved in seawater ($p$CO$_2$) increased by 1.2–2.8 µatm per year for time series from several featured ocean stations.

Overview

This tenth WMO/GAW Annual GHG Bulletin reports atmospheric abundances and rates of change of the most important long-lived greenhouse gases – carbon dioxide, methane and nitrous oxide – and provides a summary of the contributions of the other gases. These three, together with CFC-12 and CFC-11, account for approximately 96% of radiative forcing due to LLGHGs (Figure 1). For the first time, this bulletin contains a section on ocean acidification prepared in collaboration with the International Ocean Carbon Coordination Project (IOCCP) of the Intergovernmental Oceanographic Commission of the United Nations Educational, Scientific and Cultural Organization (IOC-UNESCO), the Scientific Committee on Oceanic Research (SCOR), and the Ocean Acidification International Coordination Centre (OA-ICC) of the International Atomic Energy Agency (IAEA).

The WMO Global Atmosphere Watch Programme (http://www.wmo.int/gaw) coordinates systematic observations and analysis of atmospheric greenhouse gases and other trace species. Sites where greenhouse gases are monitored in the last decade are shown in Figure 2. This map includes several featured stations where observations of CO$_2$ in ocean water are performed. Atmospheric measurement data are reported by participating institutions and archived and distributed by the World Data Centre for Greenhouse Gases (WDCGG) at the Japan Meteorological Agency.

Table 1. Global annual mean abundances (2013) and trends of key greenhouse gases from the WMO/GAW global greenhouse gas monitoring network. Units are dry-air mole fractions, and uncertainties are 68% confidence limits.

<table>
<thead>
<tr>
<th>Gas</th>
<th>CO$_2$</th>
<th>CH$_4$</th>
<th>N$_2$O</th>
</tr>
</thead>
<tbody>
<tr>
<td>Global abundance in 2013$^4$</td>
<td>396.0±0.1 ppm</td>
<td>1824±2 ppb</td>
<td>325.9±0.1 ppb</td>
</tr>
<tr>
<td>2013 abundance relative to year 1750$^a$</td>
<td>142%</td>
<td>253%</td>
<td>121%</td>
</tr>
<tr>
<td>2012–2013 absolute increase</td>
<td>2.9 ppm</td>
<td>6 ppb</td>
<td>0.8 ppb</td>
</tr>
<tr>
<td>2012–2013 relative increase</td>
<td>0.74%</td>
<td>0.33%</td>
<td>0.25%</td>
</tr>
<tr>
<td>Mean annual absolute increase during last 10 years</td>
<td>2.07 ppm/yr</td>
<td>3.8 ppb/yr</td>
<td>0.82 ppb/yr</td>
</tr>
</tbody>
</table>

$^a$ Assuming a pre-industrial mole fraction of 278 ppm for CO$_2$, 722 ppb for CH$_4$ and 270 ppb for N$_2$O. Stations used for the analyses numbered 124 for CO$_2$, 121 for CH$_4$ and 33 for N$_2$O. (An updated pre-industrial CH$_4$ value was used compared with previous bulletins, which slightly reduces radiative forcing and the relative increase in CH$_4$ relative to 1750. It had no impact on the AGGI.)
Figure 3. Globally averaged CO\textsubscript{2} mole fraction (a) and its growth rate (b) from 1984 to 2013. Differences in successive annual means are shown as shaded columns in (b).

Figure 4. Globally averaged CH\textsubscript{4} mole fraction (a) and its growth rate (b) from 1984 to 2013. Differences in successive annual means are shown as shaded columns in (b).

Figure 5. Globally averaged N\textsubscript{2}O mole fraction (a) and its growth rate (b) from 1984 to 2013. Differences in successive annual means are shown as shaded columns in (b).

Table 1 provides globally averaged atmospheric abundances of the three major LLGHGs in 2013 and changes in their abundances since 2012 and 1750. The results are obtained from an analysis of datasets (WMO, 2009) that are traceable to WMO World Reference Standards. Data from mobile stations, with the exception of NOAA sampling onboard ships transecting the Pacific Ocean, are not used for this global analysis.

The three greenhouse gases shown in Table 1 are closely linked to anthropogenic activities and they also interact strongly with the biosphere and the oceans. Predicting the evolution of the atmospheric content of greenhouse gases requires quantitative understanding of their many sources, sinks and chemical transformations in the atmosphere. Observations from GAW provide invaluable constraints on the budgets of these and other LLGHGs, and they are used to verify emission inventories and evaluate satellite retrievals of LLGHG column averages.

The NOAA Annual Greenhouse Gas Index in 2013 was 1.34, representing a 34% increase in total radiative forcing (relative to 1750) by all LLGHGs since 1990 and a 1.5% increase from 2012 to 2013 (Figure 1). The total radiative forcing by all LLGHGs in 2013 corresponds to a CO\textsubscript{2} equivalent mole fraction of 479 ppm. (http://www.esrl.noaa.gov/gmd/aggi).

Carbon dioxide (CO\textsubscript{2})

Carbon dioxide is the single most important anthropogenic greenhouse gas in the atmosphere, contributing ~65%\cite{5} to radiative forcing by LLGHGs. It is responsible for ~84% of the increase in radiative forcing over the past decade and ~83% over the past five years. The pre-industrial level of ~278 ppm represented a balance of relatively large annual two-way fluxes between the atmosphere and oceans (~80 PgC yr\textsuperscript{-1}) and the atmosphere and terrestrial biosphere (~120 PgC yr\textsuperscript{-1}). Atmospheric CO\textsubscript{2} reached 142% of the pre-industrial level in 2013, primarily because of emissions from combustion of fossil fuels and cement production (CO\textsubscript{2} emissions were 9.7±0.5 PgC\cite{6} in 2012, according to http://www.globalcarbonproject.org). This conclusion is consistent with GAW measurements of the spatial distribution of CO\textsubscript{2} at the Earth’s surface and its rate of increase, a decrease in the abundance of atmospheric oxygen (O\textsubscript{2}), and a decrease in carbon isotope ratio, 13C/12C, in atmospheric CO\textsubscript{2}. Minor contributions to increased CO\textsubscript{2} come from deforestation and other land-use change (1.0±0.5 PgC in 2012), although the net effect of terrestrial biosphere fluxes is as a sink. The average increase in atmospheric CO\textsubscript{2} from 2003 to 2013 corresponds to ~45% of the CO\textsubscript{2} emitted by human activity with the remaining ~55% removed by the oceans and the terrestrial biosphere. The main sinks for CO\textsubscript{2} emissions from fossil fuel combustion are the oceans and terrestrial biosphere. Knowledge of partitioning between these sinks is based on GAW observations of atmospheric CO\textsubscript{2} and O\textsubscript{2}. Uptake of atmospheric CO\textsubscript{2} by the oceans results in ocean acidification (see next page).

Globally averaged CO\textsubscript{2} in 2013 was 396.0±0.1 ppm (Figure 3 (a)). The increase in global annual mean CO\textsubscript{2} from 2012 to 2013 of 2.9 ppm is greater than the increase from 2011 to 2012, the average growth rate for the 1990s (~1.5 ppm yr\textsuperscript{-1}), and the average growth rate for the past decade (~2.1 ppm yr\textsuperscript{-1}). Recent increases in emissions of CO\textsubscript{2} from fossil fuel combustion (~2% yr\textsuperscript{-1} or ~0.2 PgC yr\textsuperscript{-1}) cannot explain the interannual variability in CO\textsubscript{2} growth rate nor the greater-than-average increase in annual means from 2012 to 2013. Measurements of 13C/12C in atmospheric CO\textsubscript{2} by GAW participants indicate that changes in CO\textsubscript{2} growth rate result
The ocean currently absorbs one fourth of anthropogenic CO₂ emissions, reducing the increase in atmospheric CO₂ that would otherwise occur because of fossil fuel combustion. Enhanced ocean CO₂ uptake alters the marine carbonate system that controls seawater acidity. As CO₂ dissolves in seawater it forms carbonic acid (H₂CO₃), a weak acid that dissociates into bicarbonate (HCO₃⁻) and hydrogen ions (H⁺). Increased H⁺ means increased acidity (lower pH). The ocean’s acidity increase is already measurable as oceans take up ~4 kg CO₂ per day per person. The rate of acidification is limited by the presence of carbonate ion (CO₃²⁻), which binds up most of the newly formed H⁺, forming bicarbonate. Yet that buffering reaction consumes CO₂, reducing the chemical capacity of the near-surface ocean to take up more CO₂. Currently that capacity is only 70% of what it was at the beginning of the industrial era, and it may well be reduced to only 20% by the end of the twenty-first century. The current rate of ocean acidification appears unprecedented at least over the last 300 million years, based on proxy-data from paleo archives. Acidification will continue to accelerate at least until mid-century, based on projections from Earth system models. Acidification rates are slightly affected by climate change, but those effects amount to less than 10% of the changes due to increasing CO₂. Yet freshening, for example from enhanced ice melt in the Arctic, can significantly accelerate acidification rates.

Whereas the chemistry of ocean acidification is generally well understood from observations and models, the potential consequences of ocean acidification on marine organisms are inherently more complicated. A major concern is the response of calcifying organisms, such as corals, coralline algae, mollusks and some plankton, because their ability to build shell or skeletal material (via calcification) depends on the abundance of CO₃²⁻. For many organisms, calcification declines with increased acidification. Other impacts of acidification on marine biota include reduced survival, development and growth rates, as well as changes in physiological functions and reduced biodiversity.

The decrease in surface ocean pH and increase in the surface partial pressure of CO₂ (pCO₂) due to ocean uptake of anthropogenic carbon is already detectable. These trends have been assessed and quantified by sustained ocean-time-series observations, which provide the long, temporally resolved datasets needed to characterize changes in ocean biogeochemistry and ecosystems. Examples of such time series are presented in Figure 6 and Table 2. At all locations (>10 years of measurements), surface pCO₂ has increased with time due to the rise of anthropogenic CO₂, accompanied by a corresponding acidification. These time-series stations generally represent large regions, from subpolar (MUNIDA, Station P and KNOT/K2) to subtropical (BATS, HOT, ESTOC, DYFAMED) and tropical (CARIACO, 137°E section). De-seasonalized surface data from these time series were used and a linear trend was fit over the range of measurements to provide comparable decadal changes for these stations. While there are seasonal and interannual variations, pCO₂ has increased at a rate of 1.2–2.8 µatm yr⁻¹ (the atmospheric mole fraction increased at an average rate of 2.1 ppm yr⁻¹ during the last 10 years), while ocean surface pH has decreased during the period of observations at an average rate of ~0.0011 yr⁻¹ to ~0.0024 yr⁻¹, depending on the location. The rate of these changes depends not only on the chemistry but also on other factors particular to each region. At ESTOC, HOT and BATS, the rates of increase in surface pCO₂ match the rise in atmospheric pCO₂. Similarly, the pH decrease has been steady, consistent with the rate of acidification expected from equilibration with the atmosphere. At 10°N in the 137°E section, the rise of pCO₂ and decrease of pH was slightly lower, potentially linked to changes in the circulation of the subtropical gyre. The DYFAMED time series has experienced circulation changes overlapping a strong seasonal signal, which leads to large error-bars on the linear trend. At CARIACO, surface pCO₂ was one of the highest measured, particularly during times of upwelling of CO₂-rich waters. Its large increase in pCO₂ and decrease in pH over time can be attributed to warming of surface waters linked to a reduction in upwelling, which leads to lower biological productivity. The subpolar time series displayed some of the highest extremes, caused in part by the large seasonal differences in temperature and biological productivity. At MUNIDA, the rate of increase of surface pCO₂ was low, whereas at KNOT/K2, pCO₂ increase was similar to the atmospheric rise; however, pH decrease at KNOT/K2 was one of the greatest. Sampling at Station P has not been as regular, but decreases in pH and increases in pCO₂ follow similar trends as at the other sites.

While the focus here is on surface observations, a large portion of the water column is being affected by anthropogenic carbon uptake and associated changes in ocean chemistry. Unfortunately, there are currently only a few time-series stations that conduct CO₂ measurements and this poses significant obstacles to quantifying trends in ocean acidification and carbonate chemistry. CO₂ observations from time-series stations, together with surface and interior ocean measurements from ships, need to be sustained and extended.
Figure 6. Time series of de-seasonalized surface seawater pH and respective trendlines (left) and of de-seasonalized surface $p\text{CO}_2$ (µatm) and respective trendlines (right). Featured time series include the Bermuda Atlantic Time-series Study (BATS; blue) the European Station for Time Series in the Ocean near the Canary Islands (ESTOC; purple), the Hawaii Ocean Time-series (HOT; grey); CARIACO (red); Station P (black); MUNIDA (green); the Kyodo North Pacific Time series (KNOT; orange); the station known as the Dynamics of Atmospheric Fluxes in the Mediterranean Sea (DYFAMED; yellow); the Japan Meteorological Agency 137°E section repeat hydrographic line at 10°N, 137°E (137°E section; pink). The locations of the featured time series are shown in Figure 2. Temporal sampling resolution varies from monthly to annually.

Table 2. Linear trends and standard errors for surface pH$^{a}$ and $p\text{CO}_2$ at the nine featured ocean time series

<table>
<thead>
<tr>
<th>Time series</th>
<th>pH (yr$^{-1}$)</th>
<th>$p\text{CO}_2$ (µatm yr$^{-1}$)</th>
<th>Reference</th>
</tr>
</thead>
<tbody>
<tr>
<td>BATS$^b$</td>
<td>$-0.0017\pm0.0001$</td>
<td>$1.75\pm0.08$</td>
<td>Bates et al., 2014</td>
</tr>
<tr>
<td>ESTOC$^b$</td>
<td>$-0.0014\pm0.0001$</td>
<td>$1.78\pm0.15$</td>
<td>Bates et al., 2014</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>Gonzalez-Davila et al., 2010</td>
</tr>
<tr>
<td>HOT$^b$</td>
<td>$-0.0017\pm0.0001$</td>
<td>$1.89\pm0.15$</td>
<td>Bates et al., 2014</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>Dore et al., 2009</td>
</tr>
<tr>
<td>CARIACO$^b$</td>
<td>$-0.0024\pm0.0003$</td>
<td>$2.79\pm0.37$</td>
<td>Bates et al., 2014</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>Astor et al., 2013</td>
</tr>
<tr>
<td>DYFAMED$^b$</td>
<td>$-0.0019\pm0.0009$</td>
<td>$2.56\pm0.85$</td>
<td>Touratier and Goyet, 2011</td>
</tr>
<tr>
<td>MUNIDA$^b$</td>
<td>$-0.0016\pm0.0003$</td>
<td>$1.55\pm0.24$</td>
<td>Bates et al., 2014</td>
</tr>
<tr>
<td></td>
<td></td>
<td></td>
<td>Currie et al., 2011</td>
</tr>
<tr>
<td>KNOT/K2$^b$</td>
<td>$-0.0024\pm0.0007$</td>
<td>$2.22\pm0.67$</td>
<td>Wakita et al., 2013</td>
</tr>
<tr>
<td>Station P$^c$</td>
<td></td>
<td>$c$</td>
<td>Wong et al., 2010</td>
</tr>
<tr>
<td>137°E section at 10°N$^d$</td>
<td>$-0.0011\pm0.0001$</td>
<td>$1.15\pm0.06$</td>
<td>Midorikawa et al., 2010</td>
</tr>
</tbody>
</table>

$^{a}$ pH in total hydrogen ion concentration scale at in situ temperatures.

$^{b}$ Data of pH and $p\text{CO}_2$ were calculated from measurements of total dissolved inorganic carbon, total alkalinity, temperature and salinity (silicate and phosphate assumed to be zero), except for those instances where direct measurements of pH were available. De-seasonalized data and linear trends were calculated according to GAW Report No. 184.

$^{c}$ Too few data are available for Station P for a linear trend calculation.

$^{d}$ Data of pH were calculated from measurements of $p\text{CO}_2$ and total dissolved inorganic carbon, temperature and salinity.
from small changes in fluxes between the atmosphere and terrestrial biosphere. Typically, ~120 PgC is exchanged between the atmosphere and terrestrial biosphere each year. This accounts for the observed seasonal cycle in atmospheric CO₂ abundance in the northern hemisphere. Small interannual variability (1–2%) in these fluxes, either from a change in the balance between photosynthesis and respiration or the amount of biomass burned, have a large impact on the growth rate of CO₂ (~4 PgC yr⁻¹). It is too early to say which factors are responsible for the larger-than-average increase in annual means from 2012 to 2013, but this active area of research relies on measurements by GAW participants.

**Methane (CH₄)**

Methane contributes ~17%[5] to radiative forcing by LLGHGs. Approximately 40% of methane is emitted into the atmosphere by natural sources (e.g. wetlands and termites), and about 60% comes from anthropogenic sources (e.g. ruminants, rice agriculture, fossil fuel exploitation, landfills and biomass burning). As a result of increased anthropogenic emissions, atmospheric CH₄ reached 253% of its pre-industrial level (~722 ppb) in 2013. Atmospheric CH₄ increased from ~1650 ppb in the early 1980s to a new high of 1824±2 ppb in 2013 (Figure 4 (a)). Its growth rate (Figure 4 (b)) decreased from ~13 ppb yr⁻¹ during the early 1980s to near zero during 1999–2006. Superimposed on top of the long-term changes in growth rate is significant interannual variability (IAV). Studies of IAV help understand the processes that contribute to CH₄ emissions and losses. Since 2007, atmospheric CH₄ has been increasing again; its global annual mean increased by 6 ppb from 2012 to 2013. Studies using GAW CH₄ measurements indicate that increased CH₄ emissions from wetlands in the tropics and from anthropogenic sources at mid-latitudes of the northern hemisphere are likely causes. As shown in WMO Greenhouse Gas Bulletin No. 9, increased emissions from the Arctic did not contribute to the continued increase in atmospheric CH₄ since 2007.

**Nitrous oxide (N₂O)**

Nitrous oxide contributes ~6%[5] to radiative forcing by LLGHGs. It is the third most important contributor to LLGHG radiative forcing and has the largest emissions of substances that deplete stratospheric ozone (O₃) when weighted by ozone-depleting potential. Prior to industrialization, the atmospheric N₂O burden reflected a balance between emissions from soils and the ocean, and chemical losses in the stratosphere. In the industrial era, additional anthropogenic emissions are from synthetic nitrogen fertilizers (direct emissions from agricultural fields and indirect emissions from waterways affected by agricultural runoff), fossil fuel combustion, biomass burning and other minor processes. Currently, anthropogenic sources emit ~40% of total emissions; that total, determined from GAW measurements of globally averaged N₂O (Figure 5 (a)) and its rate of increase in recent years (Figure 5 (b)), is about 16 TgN yr⁻¹.[7] Synthetic nitrogen fertilizers are the largest contributor to the increase since pre-industrial times. The globally averaged N₂O mole fraction in 2013 reached 325.9±0.1 ppb, which is 0.8 ppb greater than the previous year and 121% of the pre-industrial level (270 ppb). The increase in annual means from 2012 to 2013 is comparable to the mean growth rate over the past 10 years (0.82 ppb yr⁻¹).

GAW N₂O measurements have been used with atmospheric chemical transport models to estimate emissions at regional to continental spatial scales. Recent studies have identified tropical and subtropical land regions as the largest source regions (Thompson et al., 2014) and significant trends in N₂O emissions from Asia (Saikawa et al., 2014). Despite these advances in understanding the N₂O budget, improvements to inter-network compatibility of measurements by GAW participants are necessary. Because atmospheric N₂O has a long atmospheric lifetime (130 yr), spatial gradients are small. So, to infer estimates of emissions from the data using a transport model, biases among measurement programmes must be small, <0.1 ppb, a target that is difficult to reach with commonly used measurement technologies.

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**Figure 7.** Monthly mean mole fractions of sulphur hexafluoride (SF₆) and a suite of halocarbons (SF₆ and minor halocarbons (a) and major halocarbons (b)). The numbers of stations used for the global analyses are as follows: SF₆ (23), CFC-11 (24), CFC-12 (25), CFC-113 (23), CCl₃ (21), CH₃CCl₃ (23), HCFC-141b (9), HCFC-142b (13), HCFC-22 (13), HCFC-134a (9) and HFC-152a (8).
Other greenhouse gases

Sulphur hexafluoride is a potent LLGHG. Its emissions are almost entirely anthropogenic, and it is used mainly as an electrical insulator in power distribution equipment. Its current mole fraction is about twice the level observed in the mid-1990s (Figure 7 (a)). It has a very long atmospheric lifetime, 3 200 yr, so emissions accumulate in the atmosphere and can be determined from the rate of increase measured by GAW participants. An analysis of observations shows that emissions reported to the United Nations Framework Convention on Climate Change are greatly underestimated and are not consistent with atmospheric measurements (Levin et al., 2010). GAW SF₆ observations have another important use: they validate atmospheric mixing used in chemical transport models as described above for N₂O.

The stratospheric ozone-depleting chlorofluorocarbons (CFCs), together with minor halogenated gases, contribute ~12%[5] to radiative forcing by LLGHGs. While CFCs and most halons are decreasing, hydrochlorofluorocarbons (HCFCs) and hydrofluorocarbons (HFCs), which are also potent greenhouse gases, are increasing at relatively rapid rates, although they are still low in abundance (at ppt[6] levels, Figure 7 (a) and (b)).

This bulletin primarily addresses LLGHGs. Relatively short-lived tropospheric ozone has a radiative forcing comparable to that of the halocarbons. Many other pollutants, such as carbon monoxide, nitrogen oxides and volatile organic compounds, although not referred to as greenhouse gases, have small direct or indirect effects on radiative forcing. Aerosols (suspended particulate matter), too, are short-lived substances that alter the radiation budget. All gases mentioned herein, as well as aerosols, are monitored by the GAW Programme, with support from WMO Member countries and contributing networks.

Distribution of the bulletins

The WMO Secretariat prepares and distributes these bulletins in cooperation with the World Data Centre for Greenhouse Gases at the Japan Meteorological Agency and the GAW Scientific Advisory Group for Greenhouse Gases, with the assistance of the NOAA Earth System Research Laboratory (ESRL). The bulletins are available through the GAW Programme or the WDCGG web page.

Acknowledgements and links

Fifty WMO member countries have contributed CO₂ data to the GAW WDCGG. Approximately 46% of the measurement records submitted to WDCGG are obtained at sites of the NOAA ESRL cooperative air-sampling network. For other networks and stations, see GAW Report No. 206 (available at http://www.wmo.int/gaw). The Advanced Global Atmospheric Gases Experiment (AGAGE) also contributed observations to this bulletin. Furthermore, the GAW monitoring stations contributing data to this bulletin, shown in Figure 2, are included in the list of contributors on the WDCGG web page (http://ds.data.jma.go.jp/gmd/wdcgg/). They are also described in the GAW Station Information System, GAWSIS (http://gaw.empa.ch/gawsis) supported by MeteoSwiss, Switzerland.

The summary on ocean acidification and trends in ocean pCO₂ was jointly produced by the International Ocean Carbon Coordination Project of the Intergovernmental Oceanographic Commission of UNESCO, the Scientific Committee on Oceanic Research, and the Ocean Acidification International Coordination Centre of the International Atomic Energy Agency with support from WMO. Particular thanks go to Y. Astor, N. Bates, M. Church, L. Coppola, K. Currie, M. González-Dávila, L. Miller, T. Nakano and M. Wakita for their time-series data contribution and interpretation.

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Selected greenhouse gas observatories

The Australian Tropical Atmospheric Research Station (ATARS) at Gunn Point (12.25°S, 131.05°E, 25 m a.s.l.) is located near Darwin, in Australia’s Northern Territory. The station was established in 2010 and became a GAW regional station in 2012. It is co-located with a Bureau of Meteorology (BoM) research radar station and the operation is supported by both BoM and the Commonwealth Scientific and Industrial Research Organization (CSIRO, Australia). Gunn Point complements the growing atmospheric observation network in the Asian–Australian tropical region and is an important addition to the globally under-sampled tropical regions. Measurements currently include the main greenhouse and related trace gases (in situ CO$_2$, CH$_4$ and radon; a flask air sampling programme for CO$_2$, CH$_4$, N$_2$O, carbon monoxide (CO), hydrogen (H$_2$), $^{13}$CO$_2$, $^{12}$CO$_2$, CO, as well as aerosol scattering coefficient, black carbon, ozone, NO$_x$, gaseous elemental mercury and short-lived halocarbons of tropical coastal marine origin. In the near future, measurement programmes will expand to include in situ N$_2$O, CO, particle number concentrations and volatile organic compounds.

Pha Din (21.57°N 103.52°E, 1466 m a.s.l.) is a recently designated GAW regional station. It is located on a hill in north-western Viet Nam above the surrounding forests. The facility is operated by the Vietnamese National Hydrometeorological Service (NMS). The site has been a meteorological station since 2011, when a tower was built to install a future meteorological radar. The laboratory building provides air-conditioned laboratory space and also basic accommodation for station operators and guest researchers. Capacities for continuous observations of greenhouse gases and aerosol optical properties were established in early 2014 by the NMS, with support from two Swiss institutions (Empa and the Paul Scherrer Institute), through the Capacity Building and Twinning for Climate Observing Systems (CATCOS) project. CATCOS is coordinated by the Federal Office of Meteorology and Climatology MeteoSwiss and financed by the Swiss Agency for Development and Cooperation. CO$_2$, CH$_4$ and CO are measured with cavity enhanced absorption spectroscopy. The Pha Din station is the first of its kind recording greenhouse gases, surface ozone and aerosol properties in a rural setting in Viet Nam.

References

1. One per meg equals 0.0001 % change in O$_2$/N$_2$ ratio of air relative to a reference.
2. ppm = number of molecules of the gas per million molecules of dry air.
3. ppb = number of molecules of the gas per billion (10$^9$) molecules of dry air.
4. Indicated uncertainty ranges are calculated by a bootstrap method following Conway et al. (1994). This uncertainty is calculated with a confidence interval of 68% (one sigma).
5. This percentage is calculated as the relative contribution of the mentioned gas(es) to the increase in global radiative forcing caused by all long-lived greenhouse gases since 1750.
6. 1 PgC = 1 billion (10$^12$) tonnes or 10$^{15}$ g of carbon.
7. TgN = teragrams of nitrogen.
8. ppt = number of molecules of the gas per trillion (10$^{12}$) molecules of dry air.